

RESERVOIR COMPARTMENTALIZATION AND OVERPRESSURE IN THE SABLE SUBBASIN (NOVA SCOTIA) & PORCUPINE BASIN (IRELAND)



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ITRODUCTION		SU
DBJECTIVES	PROJECT RATIONALE	DE
ne objective of this study is to interpret the reservoir compartments and pressure cells of the able Subbasin and pressure data of the Porcupine Basin in order to better understand the ause(s) for the present (preserved) observed pressure distribution and gradients. The Sable Subbasin: Determine location, apparent displacement, and effect of faults on reservoir connectivity Construct stratigraphic and structural 3D models, focusing on seal and reservoir lithologies, to identify compartmentalization Create 3D models for pressure distribution and gradient, integrating reservoir architecture, fault behavior, and pore pressure data Determine mechanisms potentially responsible for overpressure formation and present distribution the Porcupine Basin: Complete a petrophysical analysis on 5 wells Integrate pressure data, and attempt to determine if there is pressure communication between vertically stacked reservoir units Create a multi-well correlation to investigate potential lateral communication across the basin	Overpressure is abnormally high subsurface pressure exceeding hydrostatic pressure at a given depth, and occurs when fluids become trapped in the pores of sedimentary rocks. Overpressure has been identified as a risk element in the Sable Sub-basin of Nova Scotia, and has been identified as a poorly understood risk element in the Porcupine Basin of Ireland. Previous work has established overpressure in the Sable Subbasin is variable in magnitude and unpredictable, not associated with specific depths or formations. Faults were assumed to be either dynamic (allow communication) or static (does not allow communication), which is inaccurate. This study has access to more recently acquired digital seismic and well log data, and the use of new software than previous studies, which allows for a novel approach for studying pressure in the region. Abnormal pressure and pressure distribution has received limited study in the Porcupine Basin. Pressure measurements have been collected (repeat formation tests, formation leak off tests, and drill stem tests) in several wells, providing a preliminary dataset to begin investigating pressure behavior in the basin. Increased understanding of the context and contributing factors to overpressure in pressure cells or compartments can reduce drilling and environmental risk during exploration and development of offshore resources.	Pres Depth

GEOLOGICAL SETTING

SABLE SUBBASIN

MIOCENE

OLIGOCENE

EOCENE

PALEOCENE

MASSTRIC.

CAMPANIAN

SANTONIAN CONIACIAN

TURONIAN

CENOMANIAN

ALBIAN

APTIAN

BARREMIAN

HAUTERVIAN

VALANGINIAN

BERRIASIAN

TITHONIAN

KIMMERIDG

OXFORDIAN

CALLOVIAN

BATHONIAN

BAJOCIAN

AALENIAN

TOARCIAN

PLIENSBACH

SINEMURIAN

HETTENGIAN

RHAETIAN

NORIAN

CARNIAN

LADINIAN

ANISIAN

The Scotian Basin is located offshore Nova Scotia; the total area of the basin is nearly 300,000 km², with half on the current continental shelf and the remaining half on the continental slope. Sedimentation into the basin has been near continuous over the past 250 Ma, and has been sourced from the Appalachian Orogen and transported by a paleo-drainage system, which included several large delta systems (e.g. Shelburne, Sable, and Laurentian). Sediments reached a maximum thickness of 18 km. The geological history of the basin represents diverse tectonic styles and an array of depositional environments including early-stage rifting, passive margin, carbonate bank, fluvial-deltaic-lacustrine, and deep water. Petroleum exploration in offshore Nova Scotia began in 1959, and (to date) includes 207 wells. Given the considerable area, this means that the basin is underexplored. The Scotian Basin comprises several sub-basins including Sable, Shelburne, Abenaki, Orpheus and Laurentian. Thick salt deposits formed from evaporation of restricted shallow marine waters, leading to the deposition of the Argo Formation. Significant sediment loading after deposition causes displacement of salt vertically and horizontally to create structures such as diapirs, pillows, and turtles. Salt structures are common along the offshore Scotian Margin, including within the Sable Subbasin, and have become a topic of increased interest due to their unique physical properties allowing for hydrocarbon reservoir preservation. Salt deformation is predominantly controlled by the rheology of the overlying sediments.

There are several energy projects active or recently active on the Scotian Margin, including Cohasset Panuke (1992-1999), Sable Offshore Energy Project (1999-present), and Deep Panuke (2013-present). Cohasset Panuke was Canada's first offshore oil project, and produced 44.5 MMbbls of oil. Peak production was in October 1993 with 37,500 bpd. The Sable Offshore Energy Project (SOEP) comprises 6 natural gas fields: Venture, South Venture, Thebaud, North Triumph, Glenelg, and Alma. There are an estimated 3 tcf of recoverable gas and 74.8 MMbbl condensate. Deep Panuke began production in the Panuke field in 2014, and has an appraised production life of 13 years but has recently been plagued with water production issues.

PLIOCENE

MIC >

BACCAR

MAG

MISAINE

MOHICAN

_ IROQUOIS & EQ.

SCATERIE

The Porcupine Basin is a north-south trending Mesozoic-Cenozoic age basin, located on the Irish continental margin. The Irish continental margin comprises a series of north-south and northeast-southwest trending basins, including the Porcupine, Rockall, Slyne, and Erris basins. These basins contain a preserved record of structural and

IBSURFACE PRESSURE

FINITION & FORMATION

ssure = density · gravity acceleration · depth



OVERBURDEN

> + + +

ROCK GRAIN

Normal Pressure

Underpressure

Pressure

-M

PORE _

Pressure is the force exerted on area, and depends on density, acceleration of gravity, and depth. Pore Pressure is the pressure of fluids within the pores of a reservoir. If impermeable lithologies (i.e. shales) form as sediments are compacted and the pore fluids are unable to escape, the fluids will then support the overburden. Reservoir pressure changes as fluids are produced, meaning any measurements should include a reference to a identifiable time (i.e. initial shut-in pressure versus final shut-in pressure). Lithostatic Pressure: pressure exerted per unit area by the overburden (also called geostatic pressure). Hydrostatic Pressure: pressure exerted per unit area by a column of water from sea level to a given depth. Overpressure: abnormally high pressure exceeding hydrostatic pressure at a given depth. It is not when a specific pressure is reached, instead it is an abnormally high amount of pressure for a particular depth. Overpressure occurs when excess pressure above hydrostatic values associated with fluids trapped in the pores of sedimentary rocks are unable to escape.





Hydrocarbon Generation: increased pressure as a result of the addition of hydrocarbon fluids



Normal Pressure

Overpressure

Pressure

Undercompaction: pressure formed due to rapid burial (typical of deltaic system) coupled with low-k sediments, where dewatering does not occur at the rate required for normal compaction, trapping fluids in pores. The weight of overlying sediments is supported partly by rock matrix and partly by pore fluid, resulting in an overpressured formation. Manifestations on well logs may appear as 'reversals' of resistivity and sonic data when overpressure is penetrated. Transgressive and regressive cycles associated with deltaic depositional environments can produce reservoir-seal pairs, which act as baffles and barriers to flow, and that contribute to

PRESSURE DISTRIBUTION PATTERNS

Normal or Underpressure

DRAINAGE PATH OPEI

SHALE WIRELINE INDICATORS

Tiered pressure systems are observed in dynamic basins, and basins with similar depositional High overpressure can produce indicators in wireline logs, specifically sonic velocity, resistivity, and histories are more likely to have comparable distribution patterns. Abnormal pressure in subsurface density measurements. The data will "reverse" and decrease below normal trend values. Resistivity

- can lead to multiple potential pressure distribution patterns:
- (I) Recessed-Tiered: underpressure zone bounded by normal pressure.

Constant

- (II) Ledged-Tiered: overpressure zone bordered by normal pressure.
- (III) Stepped-Tiered: basin containing several low-permeability units isolating hydrocarbonbearing reservoirs.

Recessed-Tiered Pressure System Ledged-Tiered Pressure System Stepped-Tiered Pressure System

Overpressure

Normal Pressure

Normal Pressure

Pressure

can reverse for reasons other then pore pressure including changes in temperature and salinity. It is important to complete temperature corrections to logs before using reversals as potential indicators for overpressure.



stratigraphic episodes that occurred previous-to and during the formation of the North Atlantic Ocean. The Porcupine Basin comprises sediments from the Carboniferous to Holocene (359.2 mya – present) in waters depth of 350 – 2000 m.

Jurassic, and Early Cretaceous that affected the basin formation and deposition of sediment. Prior to onset of rifting, the region contained a Proterozoic-age crystalline basement overlain by Carboniferous fluvial-deltaic sediments. The first rift episode, during the Triassic, resulted in the formation of the basins of the Irish continental margin through the development of extensional faults and the associated half-graben structures. The second rift episode, during the Late Jurassic, was the most important episode of the three as it resulted in the formation of extensional faults with related hanging wall basins and footwall high. These structures provided accommodation for thick, syn-rift sediments (potential future hydrocarbon reservoirs). The third rift episode, during the Early

This episode reactivated faults from the second and first rift episodes, and led to the deposition and preservation of a thick (up to 10 km) Aptian-Albian sediment succession.

200 km



temperatures, and diagenesis.

METHODS & TECHNIQUES

PHASE 1: Petrophysical Analysis & Data Compilation

- Compile data and digitize any required files
- Petrophysical analysis of wells (φ, V_{sh}, S_w, S_g, S_o, R_w, etc.) • Quality check and review pressure data
- Pressure-depth plots
- Synthetic seismograms & seismic-well ties

PHASE 2: Geocellular Modelling

- Locate and model faults
- Identify and model reservoir & seal intervals (horizons)
- Pillar grid and develop geometric model
- Develop velocity model using horizons
- **PHASE 3: Pressure Modelling**
 - Determine HWC
 - Complete property model utilizing pressure data Establish pressure cell distribution
 - Compare adjacent pressure cells for communication
 - Compile results into a fault & pressure map



62/7-1 Arcadia J-16 Basement Citnalta I-59 35/19-1 __0 km Emma N-03 43/13-1 34/19-1 -2 km Olympia A-12 Penobscot B-41 35/2-1 –4 km Penobscot L-30 Sable Island C-67 -6 km South Desbarres O-76 -8 km South Sable B-44 –10 km South Venture 1 P-60 South Venture 2 P-60 –12 km South Venture 3 P-60 –14 km South Venture O-59 Uniacke G-72 –16 km Venture 1 O-32 –18 km Venture 3 O-32 Venture 4 O-32 Venture 5 O-32 Venture B-13 Venture B-43 Venture B-52 Venture D-23 Venture H-22 West Olympia O-51 West Venture C-62

Sable Subbasin Wells Porcupine Basin Wells

West Venture N-01

West Venture N-91



PRELIMENARY RESULTS

There were three significant rift episodes as a result of crustal extension, which occurred during the Triassic, Late \bigtriangledown \bigtriangledown \sim $\sim\sim$ Cretaceous, initiated by northwest-southeast and west-east crustal extension. \sim Ś **UPPER CHALK** WYANDOT PLENUS MARI ARMORA DAWSON CANYON CENOMANIAN LIMESTONE 🗩 MARMORA SABLE **GREENSAND FORMATION** SHORTLAND LOGAN SHALE CANYON NASPAKI SHANNON GROUP MISSISAUGA \bigtriangledown UPPER KIMMERIDGE CLAY FORMATION VERRILL

CANYON

LOWER KIMMERIDGE CLAY FORMATION **KENMARE FORMATION BANTRY GROUP** NPB SHALE UNIT NO DATA Depth to Basement 12 km 10 km 43/1

8 km LIAS GROUP 6 km 4 km MERCIA MUDSTONE GROUP 2 km 0 km SHERWOOD SANDSTONE GROUP

Basin lithostratigraphy

Salt

 Bathymetry Study Wells Atlantic Oc N 0 50 100 200 Land Dolomite Sandstone Shale imestone Basalt Kilometers & Chalk Scotian Basin and Porcupine

Depth to basement in present arrangement the (top) Sable Subbasin and (bottom) Porcupine Basin.

P_{ng} hydrostatic pore pressure gradient

a_m - ratio of slopes of loading:unloading R_n- shale resistivity at hydrostatic pressure n 1.2 - drilling exponent Δt_n - sonic transit time in shale at normal P Δt - sonic transit time in shale from well logging V_p - compressional velocity at given depth V_{ml} - compressional velocity at mudline A - based on offset velocity vs effective stress B - based on offset velocity vs effective stress λ - rate of increase in velocity with effective stress

OBG - overburden stress gradient

Onset of

Overpressure

Reverse Trend

R - shale resistivity from well logging

 σ_v - overburden stress





where sand bodies are isolated by mudstone layers. The mudstone tends to be thinner and discontinuous, providing a limited seal and allowing for partial communication between the sand bodies – similar to a pressure-relief valve allowing for a slow release. There is also considerable variability in whether the top overpressure is "hard" or "soft", Although the drilling mud record (DMR) or mud loggers record (MLR) are not direct indicators of pressure, they are valuable in providing insight as to what pressure were expected to be encountered. The data are also helpful in wells where no pressure measurements were recorded as they still allow some understanding on what the pore pressure would have potentially been as the mud weight cannot be too far below or over this point. The mud weights act as

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